



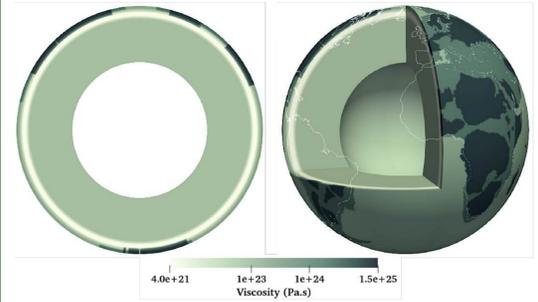
## INTRODUCTION

- Supercontinents are one of the largest spatial & temporal processes on Earth
- Contribution of slab pull to plume push mechanisms of supercontinent breakup are often discussed
- We explore the interaction between slabs & plumes beneath supercontinent in response to different lithosphere structure
- Relative contribution of plume push may be affected by continental lithosphere thickness & viscosity

## METHODS

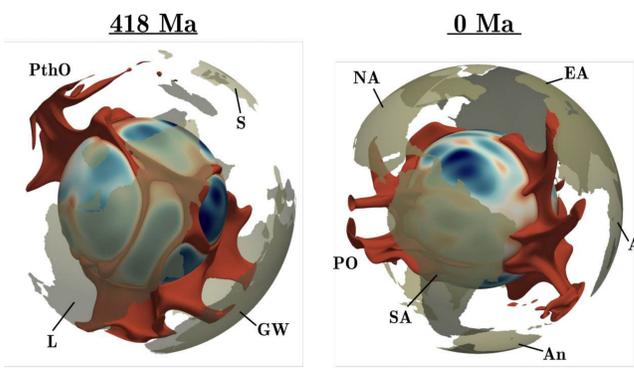
- Use 3D mantle circulation code, TERRA<sup>[1][2][3]</sup>, to simulate supercontinent cycle from 500 Ma - present
- Apply plate motion history<sup>[4]</sup>, with parameter to define continents, & cratons
- Vary thickness of continental lithosphere between 90-180 km, and vary viscosity between  $1 \times 10^{23}$  -  $1 \times 10^{25}$  Pa s

Scan QR code for input parameters & investigation parameter space

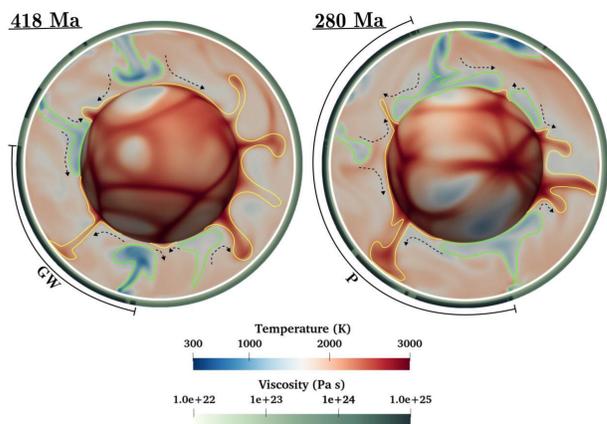


Map: Points on TERRA grid coloured by oceans, continents & cratons  
3D: Cross section and spherical visualisation of viscosity field for model 011

## PLUMES BENEATH ASSEMBLING SUPERCONTINENT



3D visualisation of reference model showing plume clusters. PthO - Panthalassic Ocean, S - Siberia, L - Laurentia, GW - Gondwana, NA - North America, EA - Eurasia, PO - Pacific Ocean, SA - South America, An - Antarctica, A - Africa

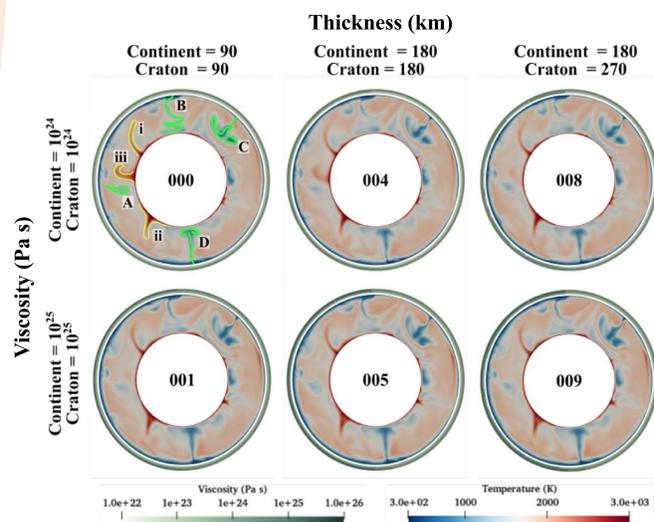


Slice through model with 180km thick continental lithosphere (viscosity = x100 oceanic lithosphere), showing the interaction between slabs (green outline) and plumes (yellow outline). GW = Gondwana, P = Pangaea

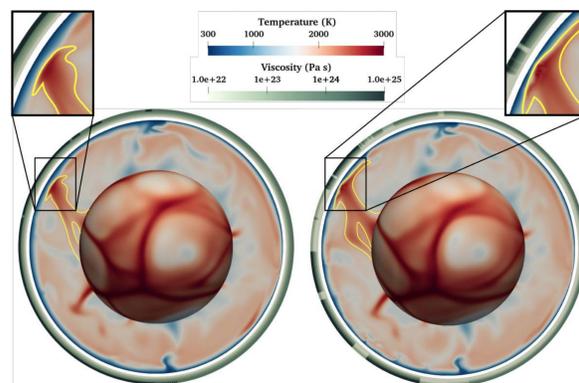
- Plumes develop into 2 antipodal clusters
- At 480 Ma, clusters beneath Gondwana & Panthalassic Ocean
- Plumes are mobile across simulations
- At 0 Ma, clusters beneath Africa & Pacific (related to LLSVPs?)

- Plumes initially develop as broad upwellings near CMB
- Downwelling slabs sweep hot upwelling material towards one central locus
- Plumes are swept beneath supercontinent by exterior subduction zones

## PLUME-SLAB INTERACTIONS WITH VARYING LITHOSPHERE STRUCTURE



2D slices through selected models at 280 Ma showing the temperature field and the viscosity field from 300-0 km depth. Left hemisphere = Pangaea, right hemisphere = Tethys Ocean



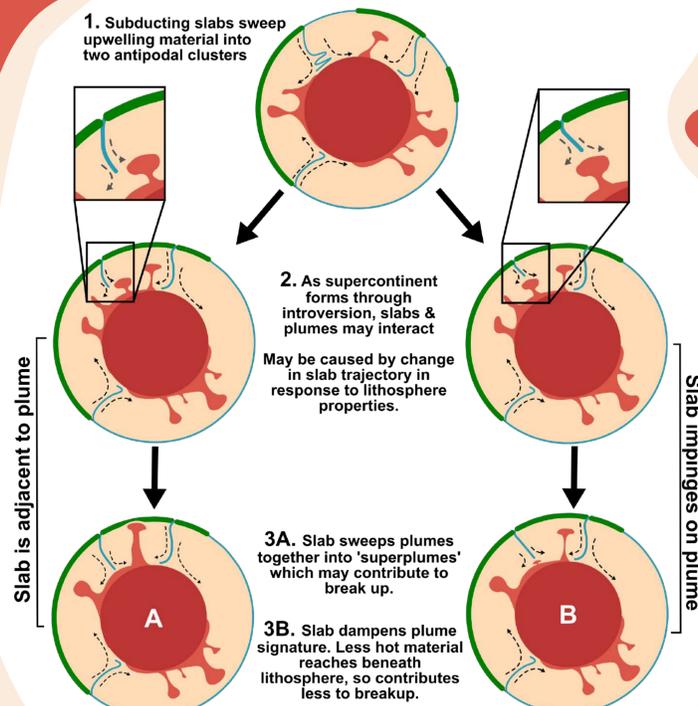
Evolution of 'superplume' at 177 Ma for model viscous continents and cratons (left) and weak, thick continents (right). Insets highlight the degree of thinning of the thermal lithosphere

- Exterior slabs (B & D) push plumes beneath supercontinent
- Slab A and plume iii may interact as interior ocean closes
- When lithosphere is viscous (or very thick), slab sinks directly above evolving plume iii
- Plume signature is dampened out in these cases
- When plume is not dampened out, it coalesces with plume i
- Larger plume thins the thermal lithosphere at time of supercontinent breakup
- Smaller plume (viscous continent models) cause less thinning and have smaller lateral extent

## DISCUSSION

- Our models have demonstrated the close relationship between slabs and plumes in the mantle
- Any large continental landmass bound by subduction zones develops a sub-continental plume, regardless of lithosphere structure
- Lifespan and extent of evolving plume varies depending on it's proximity to descending slab
- Lithosphere thickness & viscosity are two factors which may affect slab dynamics, and therefore the interaction between slabs and plumes
- Lithosphere structure during supercontinent assembly may affect the contribution of plume push forces during breakup

## CONCLUSIONS



## ACKNOWLEDGEMENTS

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